

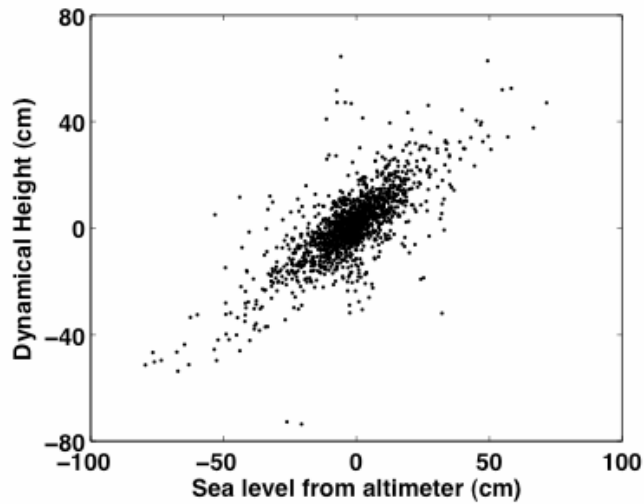
Predicting Mesoscale Variability of the North Atlantic Using a Simple, Physically-Motivated Assimilation Scheme

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Background

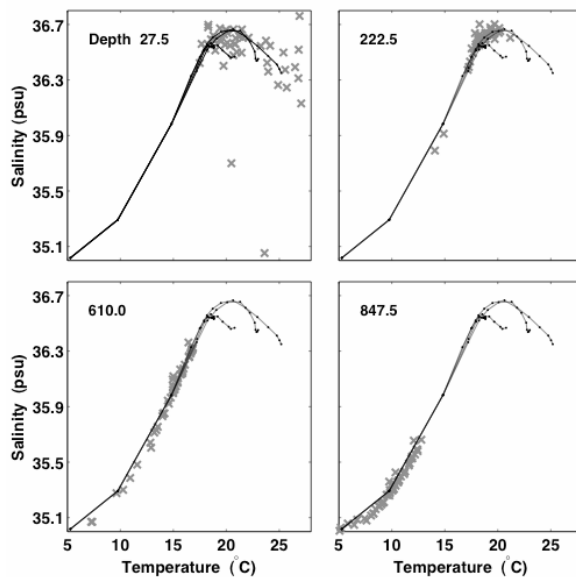
- Require computationally efficient scheme to assimilate Argo and altimeter data into regional ocean and global coupled models.
- Scales of interest are 1-30d, 10-10,000km.
- Must deal with bias in our ocean models, and rudimentary knowledge of background errors.

Observed Sea Level and Dynamic Height From Argo



- NW Atlantic
- Anomalies
- Colocated data
- RL at 1160m
- Correlation is 0.75, slope close to 1
- Simple physical balance

Argo Temperature and Salinity



- Scatterplots of T and S at different depths
- ~55.2W, 38.4N
- Complex, depth dependent structure
- Lines show Yashayaev climatology
- Shows importance of vertical advection in at depth

Modelling Uncertainty in the Background State

Motivated by these physical balances, assume

$$T = T_b - \partial T_b / \partial z \xi_D + \xi_T$$

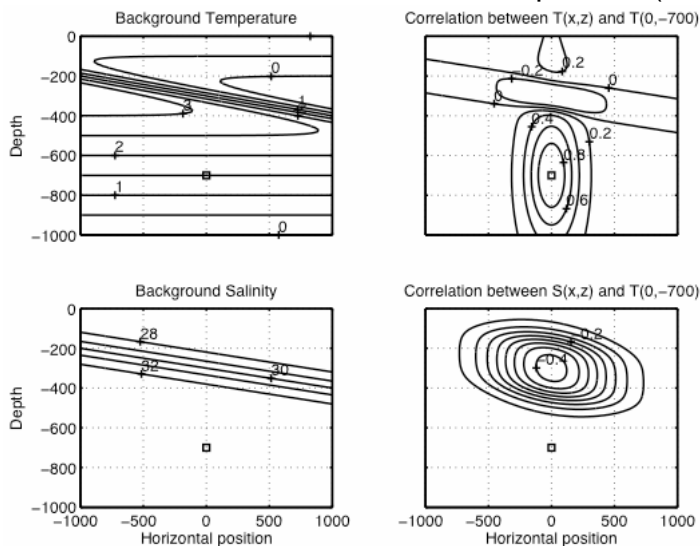
$$S = S_b - \partial S_b / \partial z \xi_D + \xi_S$$

$$\eta = \eta_b + \Delta_\rho \xi_D + \int (\alpha_T \xi_T + \alpha_S \xi_S) dz$$

Builds on: Cooper and Haines (1996), Troccoli and Haines (1999), Haines et al. (2006), Ricci et al. (2005), Weaver et al. (2006)

Implications for the B Matrix

Assume ξ_i are uncorrelated with separable (x,z) covariance:



•Complex T, S covariance

•Depends on background

State and Parameter Estimation

Let x and y be true ocean state and observation vectors, and θ a vector of uncertain parameters of covariance of ξ . Posterior pdf of state and parameters given observations is

$$p(x, \theta | y) \propto p(y | x)p(x | \theta)p(\theta)$$

Under Gaussian error assumption, maximizing posterior pdf is the same as minimizing

$$L(x, \theta) \propto \log | B_{\xi\xi}(\theta) | + J(x, \theta) - 2 \log p(\theta)$$

$$J \equiv (y - Hx)^T R^{-1} (y - Hx) + (x - x_b)^T B^{-1} (x - x_b)$$

$$B \equiv M B_{\xi\xi}(\theta) M^T$$

Online estimation of θ similar in principle to Dee (1995).

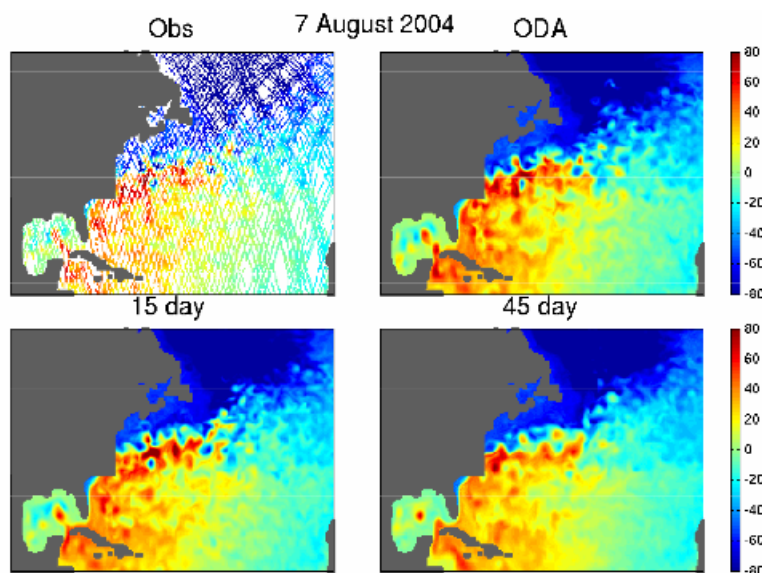
Some Details of the Scheme

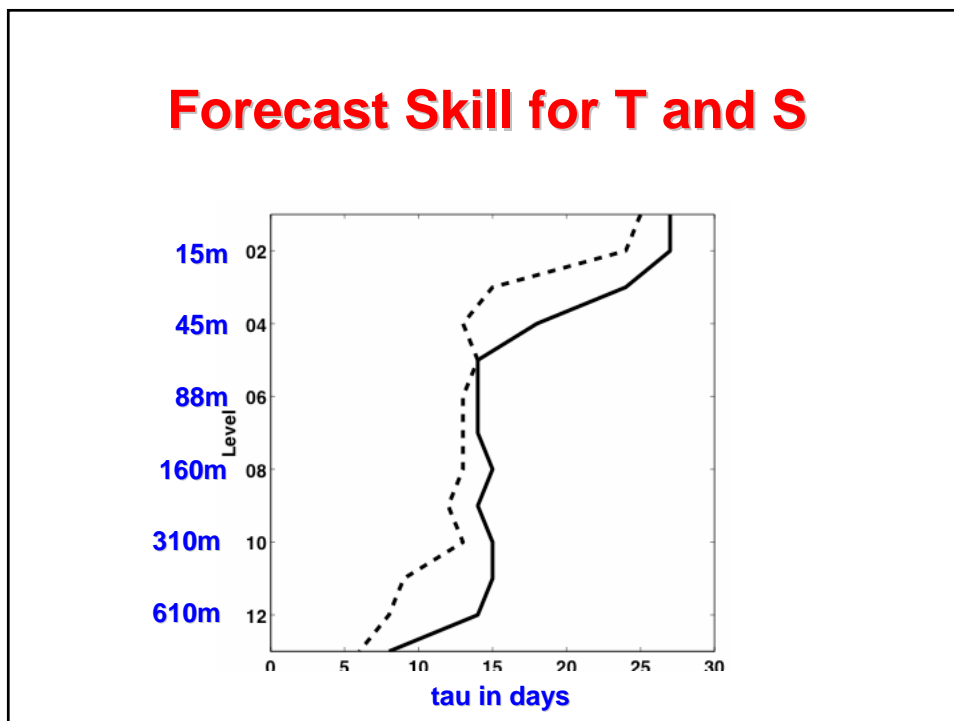
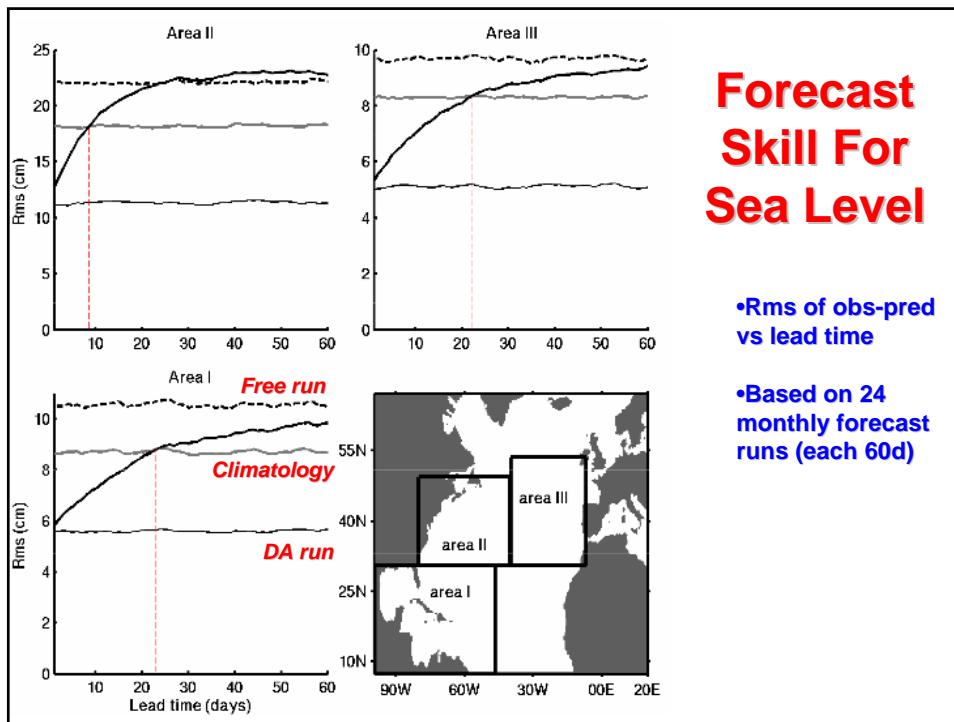
- Spectral nudging is used to suppress bias in T and S. Online, cheap and necessary. Ensures model and observed climatology match (e.g. western boundary currents separate as observed, mean PE correct) and statistics of variability are reasonable.
- Lagrangian interpolation of Argo data to analysis time.
- Estimate ocean state daily, and covariance parameters every 2 days. B changes with time (via θ) and background state (via the transformation).

North Atlantic Example

- POP ocean model, 1/3 degree, 23 levels.
- Spectrally nudged to Yashayaev monthly climatology.
- Daily atmospheric forcing from NCEP reanalysis.
- Assimilate Argo and altimeter data, 2003-5.
- Vertical gradient of background is linear combination of climatology and forecast.
- Uncertain covariance parameters (theta) are horizontal length scales and variance of the xi variables.

Typical Snapshot of Sea Level



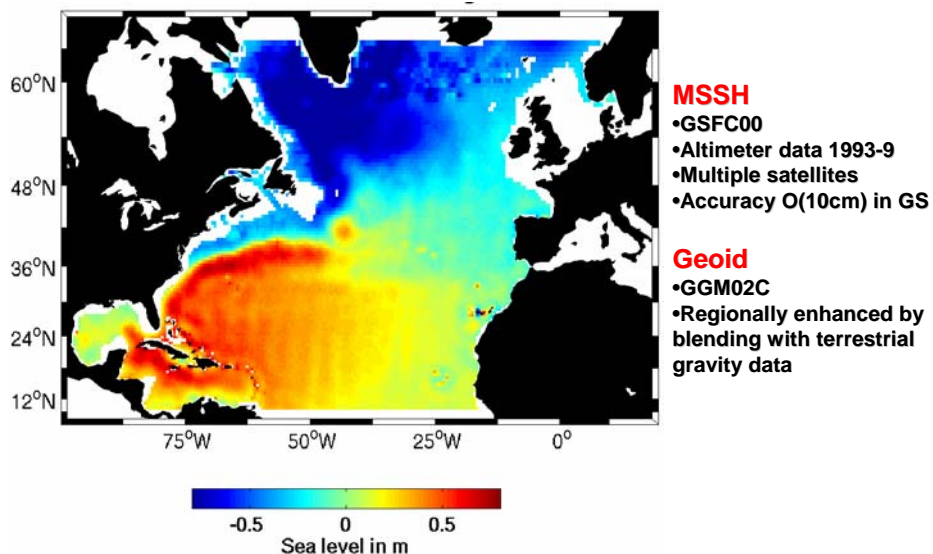


Summary

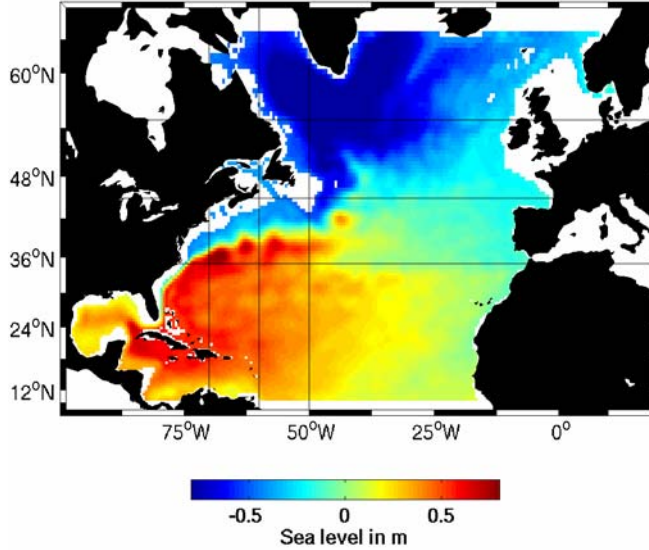
- ✓ New scheme is computationally efficient (adds 30% to run time and memory) and has useful skill.
- ✓ Plan to compare pre-operational version (with XBT) to existing operational forecast systems (e.g. SEEK).
- Works because simple physical balances built into B (which changes with time and background).
- Online bias correction is critical for forecasting North Atlantic mesoscale.
- Online estimation of B parameters gives scheme robustness and flexibility. Possible because joint posterior pdf maximized rather than marginal.

Mean Surface Topography From Space

(Jianliang Huang, NRCAN)



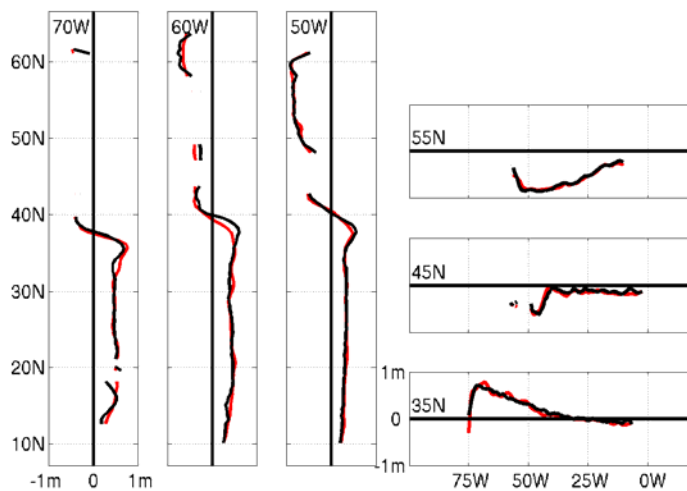
MSST of the Spectrally Nudged Model



- POP
- Realistic surface fluxes
- 1991-1999
- Spectral nudged
- Yashayaev climatology

Similar to classical diagnostic calculation based of decades of hydrographic data.

Comparing GRACE and Model-Based MSST



GRACE: Black
Ocean model: Red

*Over whole NA,
rms(error)=7.6cm!*

Spectral Nudging Gives Realistic Sea Level Variability

